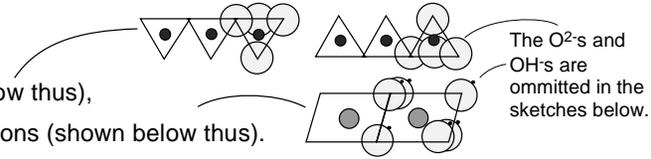


Clay Mineralogy I: Phyllosilicate minerals

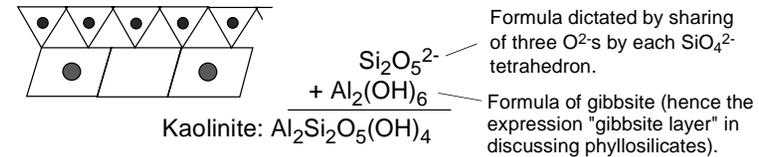
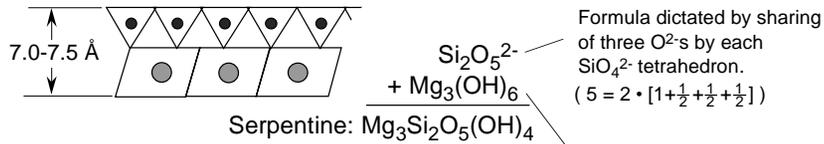
Most clay minerals can be viewed as combinations of layers of tetrahedral (T) cation sites occupied by Si^{4+} and lesser Al^{3+} or Fe^{3+} (shown below thus), and octahedral (O) cation sites occupied by Al^{3+} , Fe^{3+} , Mg^{2+} , Fe^{2+} , or other cations (shown below thus).



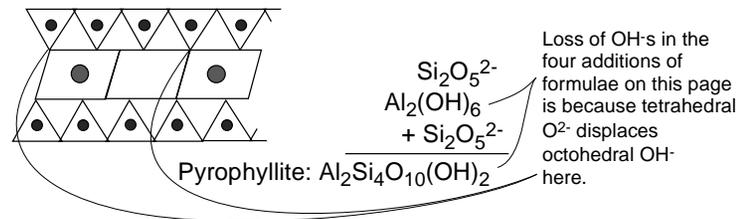
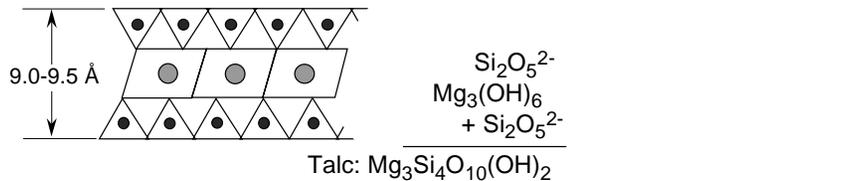
Trioctohedral minerals are those in which three of three octohedral sites are filled with 2+ cations such as Mg^{2+} or Fe^{2+} for a net charge of 6+. "Trioctohedral" thus generally means " Mg^{2+} -and/or- Fe^{2+} -bearing".

Diocctohedral minerals are those in which two of three octohedral sites are filled with 3+ cations such as Al^{3+} or Fe^{3+} for a net charge of 6+. "Diocctohedral" thus generally means " Al^{3+} -and/or- Fe^{3+} -bearing".

T-O minerals:



T-O-T minerals:



In T-O-T structures, substitution of less charged cations in the tetrahedral layers (Al^{3+} or Fe^{3+} for Si^{4+}) or in the octohedral layers (Mg^{2+} for Al^{3+} or Li^{1+} for Mg^{2+}) results in a net negative charge on the combined T-O-T layer. That allows an interlayer cation to reside between T-O-T layers. See "Clay Minerals II: T-O-T phyllosilicate minerals".

Highly generalized T-O-T + interlayer structure:

