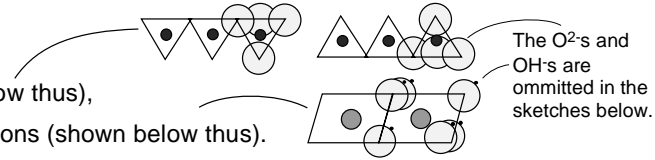


## Clay Mineralogy I: Phyllosilicate minerals

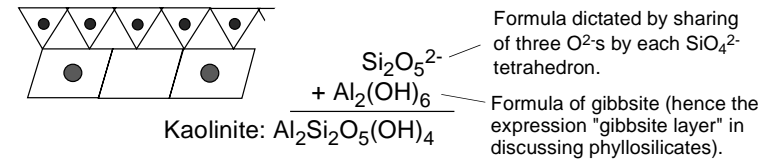
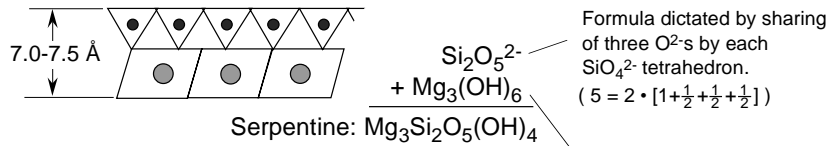
Most clay minerals can be viewed as combinations of layers of tetrahedral (T) cation sites occupied by  $\text{Si}^{4+}$  and lesser  $\text{Al}^{3+}$  or  $\text{Fe}^{3+}$  (shown below thus), and octahedral (O) cation sites occupied by  $\text{Al}^{3+}$ ,  $\text{Fe}^{3+}$ ,  $\text{Mg}^{2+}$ ,  $\text{Fe}^{2+}$ , or other cations (shown below thus).



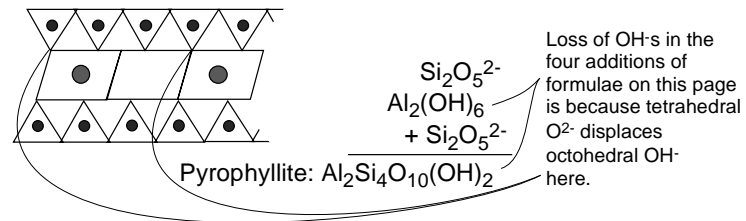
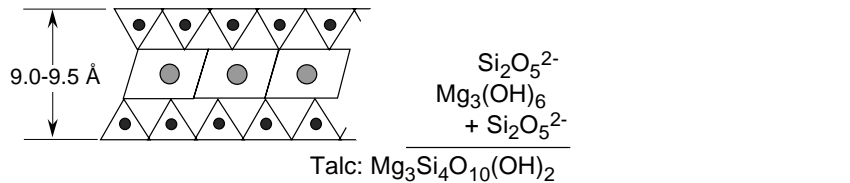
**Trioctohedral** minerals are those in which three of three octohedral sites are filled with 2+ cations such as  $\text{Mg}^{2+}$  or  $\text{Fe}^{2+}$  for a net charge of 6+. "Trioctohedral" thus generally means " $\text{Mg}^{2+}$ -and/or- $\text{Fe}^{2+}$ -bearing".

**Diocetohedral** minerals are those in which two of three octohedral sites are filled with 3+ cations such as  $\text{Al}^{3+}$  or  $\text{Fe}^{3+}$  for a net charge of 6+. "Diocetohedral" thus generally means " $\text{Al}^{3+}$ -and/or- $\text{Fe}^{3+}$ -bearing".

### T-O minerals:



### T-O-T minerals:



In T-O-T structures, substitution of less charged cations in the tetrahedral layers ( $\text{Al}^{3+}$  or  $\text{Fe}^{3+}$  for  $\text{Si}^{4+}$ ) or in the octohedral layers ( $\text{Mg}^{2+}$  for  $\text{Al}^{3+}$  or  $\text{Li}^{1+}$  for  $\text{Mg}^{2+}$ ) results in a net negative charge on the combined T-O-T layer. That allows an interlayer cation to reside between T-O-T layers. See "Clay Minerals II: T-O-T phyllosilicate minerals".

Highly generalized T-O-T + interlayer structure:

