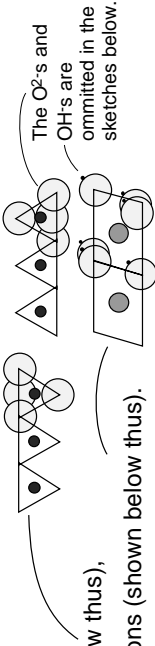


Clay Mineralogy I: Phyllosilicate minerals

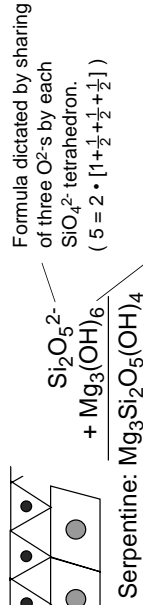
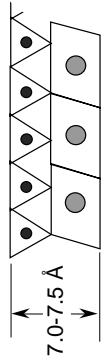
Most clay minerals can be viewed as combinations of layers of tetrahedral (T) cation sites occupied by Si^{4+} and lesser Al^{3+} or Fe^{3+} (shown below thus), and octahedral (O) cation sites occupied by Al^{3+} , Fe^{3+} , Mg^{2+} , Fe^{2+} , or other cations (shown below thus).



Trioctahedral minerals are those in which three of three octohedral sites are filled with $2+$ cations such as Mg^{2+} or Fe^{2+} for a net charge of $6+$. "Trioctahedral" thus generally means " Al^{3+} -and/or- Fe^{2+} -bearing".

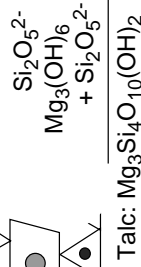
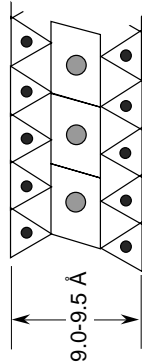
Dioctohedral minerals are those in which two of three octohedral sites are filled with $3+$ cations such as Al^{3+} or Fe^{3+} for a net charge of $6+$. "Dioctohedral" thus generally means " Al^{3+} -and/or- Fe^{3+} -bearing".

T-O minerals:



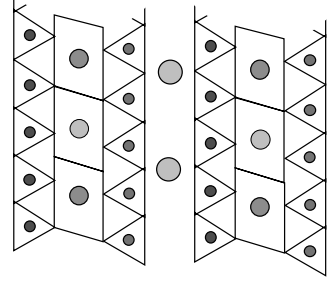
Formula of brucite (hence the expression "brucite layer" in discussing phyllosilicates).

T-O-T minerals:

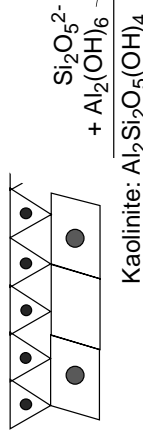


In T-O-T structures, substitution of less charged cations in the tetrahedral layers (Al^{3+} or Fe^{3+} for Si^{4+}) or in the octohedral layers (Mg^{2+} for Al^{3+} or Li^{1+} for Mg^{2+}) results in a net negative charge on the combined T-O-T layer. That allows an interlayer cation to reside between T-O-T layers. See "Clay Minerals II: T-O-T phyllosilicate minerals".

Highly generalized T-O-T + interlayer structure:



Formula dictated by sharing of three O^{2-} s by each SiO_4^{2-} tetrahedron.



Formula of gibbsite (hence the expression "gibbsite layer" in discussing phyllosilicates).

